

REMOTE SENSING BASED EVALUATION OF UNCERTAINTIES ON
MODELLING OF STREAMFLOW AFFECTED BY CLIMATE CHANGE

TAN MOU LEONG

UNIVERSITI TEKNOLOGI MALAYSIA

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TAN MOU LEONG

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DEDICATION

To my beloved mother and father

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ABSTRACT

Assessment of the impacts of land-use and climate change on streamflow is vital to develop climate adaptation strategies. However, uncertainties in the climate impact study framework could lead to changes on streamflow impact. The aim of this study is to assess the uncertainties on Digital Elevation Model (DEM), Satellite Precipitation Product (SPP) and climate projection on the modelling of streamflow affected by climate changes. These uncertainties are evaluated and reduced independently. The climate projection uncertainty is addressed through the modification of the Quantifying and Understanding the Earth System - Global Scale Impacts (QUEST-GSI) methodology. Twenty-six modified QUEST-GSI climate scenarios were used as climate inputs into the calibrated Soil and Water Assessment Tool (SWAT) model to evaluate the impacts and uncertainties of climate change on streamflow for three future periods (2015-2034, 2045-2064 and 2075-2094). The selected study areas are the Johor River Basin (JRB) and Kelantan River Basin (KRB), Malaysia. The Shuttle Radar Topography Mission (SRTM) version 4.1 (90m resolution) DEM and the Precipitation Estimation from Remotely Sensed Information using Artificial Neural Networks – Climate Data Record (PERSIANN-CDR) SPP which show a better performance were selected for the SWAT model modification, calibration and validation. The results indicated that the modified SWAT model could simulate the monthly streamflow well for both basins. Land-use and climate changes from 1985 to 2012 reduced annual streamflow of the JRB and KRB by 5% and 4.2%, respectively. In future, the annual precipitation and temperature of the JRB / KRB are projected to increase by -0.4-10.3% / 0.1-11.2% and 0.6-3.2°C / 0.8-3.3°C, respectively, and that this will lead to an increase of annual streamflow by 0.5-13.3% / 4.4-18.5%. This study showed that satellite data play an important role in providing input data to hydrological models.

ABSTRAK

Penilaian kesan guna-tanah dan perubahan iklim terhadap aliran sungai adalah penting bagi membangunkan strategi penyesuaian iklim. Namun, ketidakpastian dalam rangka kerja kajian kesan iklim akan menyebabkan perubahan kepada kesan aliran sungai. Tujuan kajian ini adalah untuk menilai ketidakpastian dalam model ketinggian berdigit (DEM), produk hujan satelit (SPP) dan unjuran iklim kepada pemodelan aliran sungai terjejas oleh perubahan iklim. Ketidakpastian ini dinilai dan dikurangkan secara berasingan. Ketidakpastian unjuran iklim ditentukan melalui pengubahsuaian metodologi hitungan dan pemahaman sistem bumi – impak skala global (QUEST-GSI). Duapuluh enam senario iklim QUEST-GSI yang telah diubahsuai digunakan sebagai input kepada model penilaian tanah dan air (SWAT) yang telah dikalibrasi untuk menilai kesan dan ketidakpastian perubahan iklim terhadap aliran sungai untuk tiga jangka masa yang akan datang (2015-2034, 2045-2064 dan 2075-2094). Tempat kajian adalah Lembangan Sungai Johor (JRB) dan Lembangan Sungai Kelantan (KRB), Malaysia. Misi topografi radar pengangkutan ulang-alik (SRTM) versi 4.1 (90m resolusi) DEM dan anggaran hujan dari informasi penderiaan jauh menggunakan jaringan neural buatan – rekod data iklim (PERSIANN-CDR) SPP yang menunjukkan prestasi lebih cermerlang telah dipilih bagi pengubahsuaian, kalibrasi dan pengesahan model SWAT. Keputusan menunjukkan model SWAT yang diubahsuai dapat mesimulasikan aliran sungai bulanan dengan baik bagi kedua-dua lembangan. Perubahan guna-tanah dan iklim dari 1985 hingga 2012 telah mengurangkan aliran sungai tahunan bagi JRB dan KRB masing-masing sebanyak 5% dan 4.2%. Hujan dan suhu tahunan masa depan bagi JRB / KRB dijangka meningkat masing-masing sebanyak -0.4-10.3% / 0.1-11.2% dan 0.6-3.2°C / 0.8-3.3°C, dan akan menyebabkan peningkatan aliran sungai tahunan sebanyak 0.5-13.3% / 4.4-18.5%. Kajian ini menunjukkan data satelit memainkan peranan penting dalam menyediakan data input kepada model hidrologi.

TABLE OF CONTENTS

CHAPTER	TITLE	PAGE
	DECLARATION	ii
	DEDICATION	iii
	ACKNOWLEDGEMENT	iv
	ABSTRACT	v
	ABSTRAK	vi
	TABLE OF CONTENTS	vii
	LIST OF TABLES	xii
	LIST OF FIGURES	xv
	LIST OF SYMBOLS	xxii
	LIST OF ABBREVIATIONS	xxvii
	LIST OF APPENDICES	xxxv
1	INTRODUCTION	1
	1.1 Background of Study	1
	1.2 Statement of the Problem	6
	1.3 Research Questions	9
	1.4 Objectives of the Study	9
	1.5 Scope of the Study	10
	1.6 Significance of the Study	11
	1.7 Structure of Thesis	13

2	LITERATURE REVIEW	14
	2.1 Introduction	14
	2.2 Hydrological Cycle	18
	2.2.1 Tropical Hydrological Cycle	20
	2.2.2 Streamflow	21
	2.3 Hydrological Modelling	22
	2.4 Remote Sensing in Hydrological Modelling	25
	2.5 Digital Elevation Model	26
	2.6 Digital Elevation Model Uncertainty Assessment	27
	2.7 Satellite Precipitation	28
	2.8 Satellite Precipitation Uncertainty Assessment	29
	2.9 Satellite Land-use	30
	2.10 Satellite Land-use Uncertainty Assessment	32
	2.11 Land-use Change and Hydrological Modelling	33
	2.12 Climate Change and Hydrological Modelling	35
	2.13 Climate Scenarios	36
	2.14 General Circulation Models	38
	2.15 Downscaling of General Circulation Models	41
	2.15.1 Hydrology Studies in Malaysia	43
	2.16 Uncertainties of Climate Change on Hydrological Modelling	44
	2.17 Climate Projection Uncertainty	45
	2.18 Studies of Climate Change Impacts on Streamflow in Asia	48
	2.19 Climate Change Impacts on Streamflow in Malaysia	53
	2.20 Summary of Literature Review	55
3	METHODOLOGY	58
	3.1 Introduction	58
	3.2 Study Area	58
	3.2.1 Johor River Basin	59
	3.2.2 Kelantan River Basin	62

3.3	Satellite Data Sets	64
3.3.1	Digital Elevation Model	65
3.3.2	Satellite Precipitation Products	70
3.3.2.1	TRMM 3B42RT and 3B42V7	71
3.3.2.2	GPCP-1DD	72
3.3.2.3	PERSIANN-CDR	72
3.3.2.4	CMORPH	72
3.3.3	Satellite Land-use	73
3.4	Observed Data Sets	74
3.4.1	Climate Data	74
3.4.2	Soil Map	76
3.4.3	Streamflow Data	79
3.5	Digital Elevation Models Uncertainty Assessment	80
3.5.1	Vertical Accuracy Assessment of Digital Elevation Models	80
3.5.2	Basin Delineation and River Network Correction	81
3.6	Satellite Precipitation Products Validation	83
3.6.1	Statistical Analysis of Satellite Precipitation Products	87
3.6.2	Bias Correction of Satellite Precipitation Product	91
3.7	Soil and Water Assessment Tool	92
3.7.1	Theoretical Concepts of Soil and Water Assessment Tool	93
3.7.2	Soil and Water Assessment Tool Modification and Set Up	97
3.7.3	Soil and Water Assessment Tool Calibration	100
3.7.4	Soil and Water Assessment Tool Validation	101
3.8	Past Land-use and Climate Changes	103
3.9	Future Climate Projection	104
3.9.1	Download General Circulation Models and Performance Assessment	104

	3.9.2 General Circulation Models Downscaling	106
	3.9.3 Modified QUEST-GSI Methodology	107
	3.10 Streamflow Analysis	107
	3.11 Summary	109
4	RESULTS AND DISCUSSION (SATELLITE DATA UNCERTAINTIES ASSESSMENT)	110
	4.1 Introduction	110
	4.2 Vertical Accuracy Assessment of Digital Elevation Models	110
	4.3 DEM-based Basin Delineation and River Network Correction	116
	4.4 Validation of Satellite Precipitation Products	126
	4.4.1 Annual Precipitation	126
	4.4.2 Seasonal Precipitation	128
	4.4.3 Monthly Precipitation	130
	4.4.4 Daily Precipitation	132
	4.5 Spatial Variability Assessment	133
	4.6 Rain Detection Ability Assessment	135
	4.7 Rain Intensity Assessment	137
	4.8 2006/2007 Flood Event Assessment	141
	4.9 Basin Assessment	143
	4.10 Hydrological Modelling Assessment	146
	4.11 Bias Correction of Satellite Precipitation Product	149
5	RESULTS AND DISCUSSION (LAND-USE AND CLIMATE CHANGES ASSESSMENT)	152
	5.1 Introduction	152
	5.2 Soil and Water Assessment Tool Parameters Sensitivity Analysis	152
	5.3 Soil and Water Assessment Tool Calibration and Validation	154
	5.4 Analysis of Historical Climate Changes	156

5.4.1	Johor River Basin	157
5.4.2	Kelantan River Basin	164
5.5	Land-use Change Detection	172
5.6	Historical Land-use and Climate Change Impacts on Streamflow	174
5.6.1	Impacts of Land-use Change (S2 vs S1)	174
5.6.2	Impacts of Climate Change (S3 vs S1)	178
5.6.3	Combined Impacts of Land-use and Climate Change (S4 vs S1)	179
5.7	Modified QUEST-GSI Methodology	180
5.7.1	General Circulation Models Performance Evaluation	180
5.7.2	Future Climate Projection	186
5.8	Impacts of Future Climate Change on Streamflow	194
5.9	Climate Projection Uncertainties Assessment	198
6	CONCLUSIONS AND RECOMMENDATIONS	204
6.1	Conclusions	204
6.1.1	Digital Elevation Model Uncertainties Assessment	205
6.1.2	Satellite Precipitation Uncertainties Assessment	205
6.1.3	Soil and Water Assessment Tool	206
6.1.4	Impacts of Historical Land-use and Climate Change on Streamflow	207
6.1.5	Impacts and Uncertainties of Future Climate Change on Streamflow	208
6.2	Recommendations for Future Work	209
	REFERENCES	211
	Appendices A-S	239-311

LIST OF TABLES

TABLE NO.	TITLE	PAGE
2.1	Popular hydrological models in climate change impact assessment	24
2.2	Satellite imagery for land use mapping	32
2.3	CMIP5 General Circulation Models	40
2.4	Advantages and disadvantages of statistical and dynamical downscaling	42
2.5	The QUEST-GSI methodology climate projection	46
2.6	Studies that employed the QUEST-GSI methodology	47
2.7	Simulated streamflow historical (1984-1993) and future periods (2025-2034 and 2041-2050) at major basins of Peninsular Malaysia	54
3.1	Total population by gender and household of Johor River Basin, Malaysia	61
3.2	Total population by gender and household of Kelantan River Basin, Malaysia	64
3.3	Basic information on SPPs used in this study	71
3.4	Details of climate stations	75
3.5	Soil information	77
3.6	Details of stream gauges	80
3.7	Thirty-eight principal rain gauges across Malaysia, from the Malaysia Meteorological Department	85
3.8	Contingency table for comparing gauge and satellite precipitation estimate.	91

3.9	Climate parameters in SWAT model	99
3.10	Soil parameters in SWAT model	100
3.11	Model performance rating guidelines	103
3.12	The modified QUEST-GSI scenarios	108
4.1	T-test statistical analysis of the GCPs elevation of topographic map and various DEMs scenarios (significance level: 0.05)	116
4.2	DEMs vertical accuracy assessment studies	117
4.3	Student t test analysis between observed and estimated precipitations	128
4.4	Statistical analysis of the annual, seasonal, monthly and daily precipitation (2003-2007) between precipitation products and rain gauges	129
4.5	Overall rain-detection capability of each precipitation products over Malaysia	136
4.6	Statistical analysis of SPPs with rain gauges for 2006/07 flood event at highly affected region from December 2006 to January 2007	142
4.7	Statistical analysis of monthly rainfall between satellite precipitation products and rain gauges over the Kelantan River Basin and Johor River Basin from 2003 to 2007	144
4.8	Statistical analysis between monthly observed and SWAT-simulated streamflow by satellite precipitation products (2003-2007) at Jam Guillemard (KRB) and Rantau Panjang (JRB) stations	149
4.9	Statistical analysis of raw and corrected PERSIANN-CDR at the Jan Guillermand (KRB) and Rantau Panjang (JRB) stations from 2003 to 2007	150
5.1	Calibrated parameters for the SWAT model in JRB (1: most sensitive)	153

5.2	Calibrated parameters for the SWAT model in KRB (1: most sensitive)	154
5.3	Analysis of Mann-Kendall statistic (Z) and Sen's slope (β) for annual and temperature precipitation in JRB	163
5.4	Analysis of Mann-Kendall statistic (Z) and Sen's slope (β) for annual precipitation and temperature in KRB	171
5.5	Land-use changes JRB (1990 vs 2008)	173
5.6	Land-use changes KRB (1990 vs 2008)	174

LIST OF FIGURES

FIGURE NO.	TITLE	PAGE
1.1	“Top down” framework for climate change impact on hydrological cycle studies	5
2.1	Schematic view of the components of the climate system, their processes and interactions	15
2.2	Mean changes of global (a) temperature and (b) precipitation over the twenty-first century according to the ensembles mean CMIP5 GCMs under RCP 8.5 scenario	16
2.3	The global hydrological storage and fluxes with anthropogenic and natural cycles	18
2.4	Tropical hydrological cycle	21
2.5	Hydrological model classification	24
2.6	Reported application of hydrological models for assessing climate change impact between 2000 and 2014	25
2.7	Workflow of climate change streamflow impact study	36
2.8	a) 2100 radiative forcing trend; (b) CO ₂ emission versus 2100 radiative forcing; and (c) greenhouse gases level of 2100 radiative forcing	38
2.9	Relative changes of (a) mean flow, (b) high flow and (c) low flow under RCP 8.5 scenario for the 2071-2100 period compared to baseline 1971-2000 period	49

2.10	Observed and projected changes in annual mean temperature and precipitation over Asia	51
2.11	Annual and seasonal streamflow changes trend in China (shaded area are dominated by decreasing streamflow)	52
2.12	Simulated historical and future annual mean air temperature of Peninsular Malaysia	55
2.13	Simulated historical and future annual mean precipitation of Peninsular Malaysia	55
3.1	Johor River Basin	60
3.2	Kelantan River Basin	63
3.3	ASTER GDEM of the Johor River Basin	66
3.4	EarthEnv-DEM90 of the Johor River Basin	67
3.5	SRTM v4.1 of the Johor River Basin	67
3.6	ASTER GDEM of the Kelantan River Basin	68
3.7	EarthEnv-DEM90 of the Kelantan River Basin	69
3.8	SRTM v4.1 of the Kelantan River Basin	70
3.9	Soil map of the Johor River Basin	78
3.10	Soil map of the Kelantan River Basin	79
3.11	River network of the Johor River Basin	82
3.12	River network of the Kelantan River Basin	83
3.13	Distribution of rain gauges (MMD = Malaysia Meteorological Department; DID = Department of Irrigation and Drainage Malaysia) and topography of Malaysia	85
3.14	Schematic representation of the hydrological cycle in SWAT	94
3.15	SWAT modelling procedure	97
4.1	GCPs well distributed over the Johor River Basin	112
4.2	GCPs well distributed over the Kelantan River Basin	113
4.3	Scatter plot of GCPs elevation of topographic map vs (a) ASTER GDEM2 30m, (b) SRTM v4.1	

	90m, and (c) EarthEnv-DEM90 90m in the Johor River Basin	114
4.4	Scatter plot of GCPs elevation of topographic map vs (a) ASTER GDEM2 30m, (b) SRTM v4.1 90m, and (c) EarthEnv-DEM90 90m in the Kelantan River Basin	115
4.5	Real, corrected and uncorrected river network for the Johor River Basin	118
4.6	Real, corrected and uncorrected river networks for region (a) one, (b) two, (c) three, (d) four, (e) five, and (f) six in the Johor River Basin	119
4.7	Real, corrected and uncorrected river network for the Kelantan River Basin	122
4.8	Real, corrected and uncorrected river networks for region (a) one, (b) two, (c) three, (d) four, (e) five, and (f) six in the Kelantan River Basin	123
4.9	Spatial distribution of average annual precipitation for the period 2003-2007 estimated from (a) 3B42RT, (b) 3B42V7, (c) GPCP-1DD, (d) PERSIANN-CDR, (e) CMORPH, and (f) rain gauges	127
4.10	Two-dimensional histogram of mean monthly precipitation between satellite precipitation products with the rain gauges for the period 2003 – 2007	131
4.11	Coefficient of determination (R^2) of daily precipitation between rain gauges and (a) 3B42RT, (b) 3B42V7, (c) GPCP-1DD, (d) PERSIANN-CDR and (e) CMORPH over Malaysia	134
4.12	Coefficient of determination (R^2) of monthly precipitation between rain gauges and (a) 3B42RT, (b) 3B42V7, (c) GPCP-1DD, (d)	

	PERSIANN-CDR and (e) CMORPH over Malaysia	135
4.13	The probability of detection (POD) of daily precipitation between satellite precipitation products and rain gauges over Malaysia	137
4.14	The occurrence probability distribution functions (PDF) of daily precipitation (2003-07) aggregated from 342 rain gauges over (a) Malaysia and (b-h) different regions of Malaysia	139
4.15	Comparison of daily precipitation series between precipitation products and selected rain gauges (highly affected area) for the 2006/2007 flood event	143
4.16	Monthly rainfall pattern between observed and satellite precipitation products at (a) station 48615, (b) 48616, (c) 48672 and (d) 48679 during 2003-2007	145
4.17	Monthly observed and SWAT-simulated streamflow by five satellite precipitation products at the (a) Jam Guillemard (KRB) and (b) Rantau Panjang (JRB) station between 2003 and 2007	148
4.18	Monthly streamflow of monthly observed against raw and corrected PERSIANN-CDR at (a) Jam Guillemard and (b) Rantau Panjang station for the period of 2003 to 2007	151
5.1	Observed and simulated streamflow at Rantau Panjang station	155
5.2	Observed and simulated streamflow at Jam Guillemard station	155
5.3	(a) Annual precipitation and (b-d) temperature changes at the 48672 station (1985-2012)	158

5.4	Monthly (a) precipitation and (b) temperature changes at the 48672 station (1985-1998 vs 1999-2012)	159
5.5	Annual (a) precipitation and (b-d) temperature changes at the 48679 station (1985-2012)	160
5.6	Annual precipitation and temperature changes at the 48679 station (1985-2012)	161
5.7	Annual precipitation trend at various (a-h) precipitation stations in JRB (1985-2012)	162
5.8	(a) annual precipitation and (b-d) temperature changes at the 48616 station (1985-2012)	165
5.9	Monthly (a) precipitation and (b) temperature changes at the 48616 station (1985-1998 vs 1999-2012)	166
5.10	Annual precipitation trend at various precipitation stations in KRB (1985-2012)	167
5.11	Land use changes of the Johor River Basin (JRB) and Kelantan River Basin (KRB) (1990 vs 2008)	173
5.12	Annual streamflow changes under land-use (S2), climate (S3) and combined impact (S4) scenarios in Johor River Basin	175
5.13	Annual streamflow changes under land-use (S2), climate (S3) and combined impact (S4) scenarios in Kelantan River Basin	175
5.14	Annual streamflow changes under land-use (S2), climate (S3) and combined impact (S4) scenarios in Johor River Basin for various (a-d) sub-basins	176
5.15	Annual streamflow changes under land-use (S2), climate (S3) and combined impact (S4) scenarios in Kelantan River Basin for various sub-basins	177
5.16	(a) NME and (b) NRMSE for annual precipitation from 1985 to 2004 in JRB. The red dotted lines are the uncertainty bounds	182

5.17	(a) NME and (b) NRMSE for annual precipitation from 1985 to 2004 in JRB. The red dotted lines are the uncertainty bounds	183
5.18	(a) NME and (b) NRMSE for annual precipitation from 1985 to 2004 in KRB. The red dotted lines are the uncertainty bounds	184
5.19	(a) NME and (b) NRMSE for annual precipitation from 1985 to 2004 in KRB. The red dotted lines are the uncertainty bounds	185
5.20	Annual (a) precipitation, (b) maximum and (c) minimum temperature changes of Ensemble_5 at Kluang station (48672) under three RCPs scenarios for the 2015-2034, 2045-2064 and 2075-2094.	188
5.21	Monthly (a) precipitation, (b) maximum and (c) minimum temperature changes of Ensemble_5 at Kluang station (48672) under three RCPs scenarios for the 2015-2034, 2045-2064 and 2075-2094	189
5.22	Annual (a) precipitation, (b) maximum and (c) minimum temperature changes of Ensemble_5 at Senai station (48679) under three RCPs scenarios for the 2015-2034, 2045-2064 and 2075-2094	190
5.23	Monthly (a) precipitation, (b) maximum and (c) minimum temperature changes of Ensemble_5 at Senai station (48679) under three RCPs scenarios for the 2015-2034, 2045-2064 and 2075-2094	191
5.24	Annual (a) precipitation, (b) maximum and (c) minimum temperature changes of Ensemble_5 at Kuala Krai station (48616) under three RCPs scenarios for the 2015-2034, 2045-2064 and 2075-2094	192

5.25	Monthly (a) precipitation, (b) maximum and (c) minimum temperature changes of Ensemble_5 at Kuala Krai station (48616) under three RCPs scenarios for the 2015-2034, 2045-2064 and 2075-2094	193
5.26	(a) Annual and (b) monthly streamflow changes at Rantau Panjang station of Ensemble_5 under RCP 2.6, RCP 4.5 and RCP 8.5 for the periods of 2020s, 2050s and 2080s	196
5.27	a) Annual and (b) monthly streamflow changes at Jam Guillemard station of Ensemble_5 under RCP 2.6, RCP 4.5 and RCP 8.5 for the periods of 2020s, 2050s and 2080s	197
5.28	Changes of annual streamflow against the baseline at Rantau Panjang station for climate scenarios: (a) Prescribed temperature, (b) GCM structure, and (c) 2°C increase in temperature	200
5.29	Changes of monthly streamflow against the baseline at Rantau Panjang station for climate scenarios: (a) Prescribed temperature, (b) GCM structure, and (c) 2°C increase in temperature	201
5.30	Changes of annual streamflow against the baseline at Jam Guillemard station for climate scenarios: (a) Prescribed temperature, (b) GCM structure, and (c) 2°C increase in temperature	202
5.31	Changes of monthly streamflow against the baseline at Jam Guillemard station for climate scenarios: (a) Prescribed temperature, (b) GCM structure, and (c) 2°C increase in temperature	203

LIST OF SYMBOLS

\bar{P}_x	- Mean of the DEM elevation
\bar{P}_{ref}	- Mean of reference elevation
$\Delta P(x, y)$	- Precipitation different between the satellite and observed data at a given point
$P_{S(x,y)}$	- Satellite precipitation at a specific location (x,y)
$P_{G(x,y)}$	- Gauge precipitation at a specific location (x,y)
$\Delta P_{(x,y)int}$	- Spatially interpolated different map
$P_{s.corrected}$	- Satellite precipitation after correction
SW_t	- Final soil water content
SW_o	- Initial soil water content on day i
R_{day}	- The amount of precipitation on day i
Q_{surf}	- Amount of surface runoff on day i
E_a	- Amount of evapotranspiration on day i
w_{seep}	- Amount of water entering the vadose zone from the soil profile on day i
Q_{gw}	- Amount of return flow on day i
Q_{surf}	- Accumulated runoff or rainfall excess
R_{day}	- Rainfall depth for the day
I_a	- Initial abstraction which includes surface storage, interception and infiltration prior to runoff
S	- Retention parameter

$f_{inf,t}$	- Infiltration rate at time t
K_e	- Effective hydraulic conductivity
ψ_{wf}	- Wetting front matric potential
$\Delta\theta_v$	- Change in volumetric moisture content across the wetting front
$F_{inf,t}$	- Cumulative infiltration at time t
E_o	- Potential evapotranspiration
λ	- Latent heat of vaporization
H_o	- Extra-terrestrial radiation
T_{mx}	- Maximum air temperature for a given day
T_{mn}	- Minimum air temperature for a given day
\bar{T}_{av}	- Mean air temperature for a given day
q_{ch}	- Rate of flow in the channel
A_{ch}	- The cross-sectional area of flow in the channel
R_{ch}	- Hydraulic radius for a given depth of flow
slp_{ch}	- Slope along the channel length
n	- Manning's coefficient for the channel
ΔV_{stored}	- Change in volume of storage during the time step
V_{in}	- Volume of inflow during the time step
V_{out}	- Volume of outflow during the time steps
q_{in}	- Inflow rate
q_{out}	- Discharge rate
K	- Storage time constant for the reach
X	- Weighting factor
\bar{o}	- Mean observed streamflow
\bar{p}	- Mean simulated streamflow
$T_{delta,daily}$	- Delta downscaled daily temperature
$T_{obs,daily}$	- Observed daily temperature
\bar{T}_{fut}	- GCM average monthly temperature of the future period

\bar{T}_{bas}	- GCM average monthly temperature of the baseline period
$P_{delta,daily}$	- Delta downscaled daily precipitation
$P_{obs,daily}$	- Observed daily precipitation
\bar{P}_{fut}	- GCM average monthly precipitation of the future period
\bar{P}_{bas}	- GCM average monthly precipitation of the baseline period
'	- Minute
"	- Second
%	- Percentage
\pm	- Plus-minus
°	- Degree
A	- Hits (event forecast to occur, and did occur)
ACC	- Accuracy
B	- False alarm (event forecast to occur, but did not occur)
C	- Misses (event forecast not to occur, but did occur)
CH_K2	- Channel effective hydraulic conductivity
CH_N2	- Manning's value for main channel
CN2	- Initial SCS Curve Number II value
CSI	- Critical success index
D	- Correct negative (event forecast not to occur but occur)
Ensemble_5	- Ensemble of five selected GCMs
ESCO	- Soil evaporation compensation factor
FAR	- False alarm ratio
G	- Gauge precipitation
GHz	- Gigahertz
GW_REVAP	- Groundwater "revap" coefficient
GWQMN	- Threshold water depth in the shallow aquifer for flow
HSS	- Heidke skill score
km	- Kilometre
km ²	- Square kilometre
km ³	- Cubic kilometre
km ³ /year	- Cubic kilometre per year
m	- Metre

m.a.s.l	- Above mean sea level
m ³ /day	- Cubic metre per day
m ³ s ⁻¹	- Meter cubic per second
ME	- Mean error
MJkg ⁻¹	- Megajoules per kilogramme
mm/day	- Millimetre per day
mm/day	- Millimeter per day
mm/month	- Milimetre per month
mm/year	- Millimetre per year
<i>n</i>	- Number of gcps
NME	- Normalize mean error
NRMSE	- Normalize root mean square error
NSE	- Nash sucliffe efficiency
<i>o</i>	- Observed streamflow
°C	- Celsius
°N	- North
°S	- South
<i>p</i>	- Simulated streamflow
PB	- Percentage bias
POD	- Probability of detection
P_{ref}	- Reference elevation (gcps from topographic map)
P_x	- DEM elevation
R ²	- Coefficient of determination
RB	- Relative bias
REVP_MN	- Threshold depth of water in the shallow aquifer for “revap” to occur
RMSE	- Root mean square error
RSR	- Ratio of root mean square error
S	- Satellite precipitation
SB	- Sub-basin
SOL_AWC	- Available water capacity
<i>t</i>	- Time
Tmax	- Maximum temperature

T _{mean}	- Mean temperature
T _{min}	- Minimum temperature
UTC	- Coordinated universal time
W _m ⁻²	- Watt per square metre

LIST OF ABBREVIATIONS

3D	- Three dimensional
ACCESS1-0	- Australia Community Climate and Earth System Simulator, version 1
AMJ	- April, May and Jun
ANION_EXCL	- Fraction of porosity (void space) from which anions are excluded
AOGCMs	- Atmospheric-ocean general circulation models
APHRODITE	- Asian Precipitation Highly Resolved Observational Data Integration Towards the Evaluation of Water Resources
AR4	- Assessment Report 4
AR5	- Assessment Report 5
ASTER	- Advanced Spaceborne Thermal Emission and Reflection Radiometer
AVHRR	- Advanced Very High Resolution Radiometer
BCC-CSM1-1	- Beijing Climate Centre Climate System Model, version 1.1
BNU-ESM	- Beijing Normal University, Earth System Model
CanESM2	- Canadian Earth System Model, version 2
Cat-PDM	- Probability Distributed Model
CCCma	- Canadian Centre for Climate Modelling and Analysis
CCSM3	- Community Climate System Model, version 3
CCSM4	- Community Climate System Model, version 4
CESM1-CAM5	- Community Earth System Model- Community Atmosphere Model
CGCM3.1	- Canadian Coupled General Climate Model, version 3.1

CGIAR-CSI	- Consortium for Spatial Information of the Consultative Group of International Agriculture Research
CLAY1	- Clay content of the first layer of soil
CMIP 3	- Coupled Model Intercomparison Project Phase 3
CMIP 5	- Coupled Model Intercomparison Project Phase 5
CMORPH	- Climate Prediction Center Morphing Technique
CN	- Curve Number
CNES	- Centre national d' études spatiales – the French space agency
CNRM-CM5	- Centre National de Recherches Meteorologiques- Coupled Global Climate Model, version 5
CRU-TS	- Climatic Research Unit- Time series
CSIRO-Mk3-6-0	- Commonwealth Scientific and Industrial Research Organization- Mark Climate Model, version 3.6
CSIRO-Mk3.0	- Commonwealth Scientific and Industrial Research Organization- Mark Climate Model, version 3.0.
DEMs	- Digital Elevation Models
DEWPT	- Average daily few point temperature per month of the year in degree celcius
DID	- Department Irrigation And Drainage
EC-EARTH	- European Centre Earth System Model
EGM96	- Earth Gravitational Model 1996
EM	- East Malaysia
EM-DAT	- International Emergency Disasters Database
ENO	- Environment and Organism
ET	- Evapotranspiration
ETM+	- Enhanced Thematic Mapper Plus
FELDA	- Federal Land Development Authority
FGOALS-g2	- Flexible Global Ocean-Atmosphere-Land System model, Grid-point version 2
FRIM	- Forest Research Institute Malaysia
GCMs	- General Circulation Models
GCPs	- Ground control points
GDA	- Geographic Differential Analysis

GDEM	- Global Digital Elevation Model
GFDL-CM3	- Geophysical Fluid Dynamics Laboratory Coupled Model, version 3
GFDL-ESM 2M	- Geophysical Fluid Dynamics Laboratory- Earth System Model, version 2
GHCN2	- modified version of Global Historical Climatology Network
GHG	- Greenhouse Gas
GLSDEM	- Global Land Survey Digital Elevation Model
GOES	- Geostationary Operational Environmental Satellite system
GPCC	- Global Precipitation Climatology Center
GPCP	- Global Precipitation Climatology Project
GPCP-1DD	- GPCP – one degree daily
GRA	- Geographical Ratio Analysis
HadCM3	- Hadley Centre Coupled Model, version 3
HadGEM1	- Hadley Centre Global Environment Model, version 1
HadGEM2-ES	- Hadley Centre Global Environment Model version 2 Earth System
HBV	- Hydrological Simulation Model
HeC-HMS	- Hydrologic Engineering Center-Hydrologic Modeling System
HESS	- Hydrology and Earth System Sciences
HRUs	- Hydrologic Response Units
HSPF	- Hydrologic Simulation Package-Fortran
HYDGRP	- Soil hydrologic group
IDW	- inverse distance weighting
INM-CM4	- Institute for Numerical Mathematics Climate Model, version 4
IPCC	- Intergovernmental on Climate Change
IPSL-CM4	- Institut Pierre Simon Laplace- Climate Model, version 4.
IPSL-CM5A-MR	- Institut Pierre Simon Laplace- Coupled Model 5A- Medium Resolution.
IR	- Thermal infrared
IRS 1D LISS-3	- Indian Remote Sensing 1D Linear Imaging Self-Scanning System III

JAS	- July, August and September
JFM	- January, February and March
JPEG	- Joint Photographic Experts Group
JRB	- Johor River Basin
JUPEM	- Jabatan Ukur dan Pemetaan Malaysia
KEJORA	- South East Johore Development Authority
KINEROS	- Kinematic Runoff and Erosion
KRB	- Kelantan River Basin
LAI	- Leaf Area Index
LAMs	- Limited Area Models
LH-OAT	- Latin hypercube sampling by one at a time design
LULC	- Land use land cover
MAE	- Mean absolute error
MARDI	- Malaysian Agriculture Research and Development Institute
ME	- Mean error
METI	- Ministry of Economy, Trade, and Industry of Japan
MIKE-SHE	- Mike- Syst ème Hydrologique Europ éen
MIROC5	- Model for Interdisciplinary Research on Climate, version 5
MIROC-ESM	- Model for Interdisciplinary Research on Climate- Earth System Models
MMD	- Malaysia Meteorological Department
MOA	- Ministry of Agriculture and Agro-based Industry Malaysia
MODIS	- Moderate Resolution Imaging Spectroradiometer
MPI	- Max Planck Institute for Meteorology
MRI-CGCM3	- Meteorological Research Institute- Coupled General Circulation Model, version 3
MSL	- Mean sea level
MW	- Microwave
NAHRIM	- National Hydraulic Research Institute Malaysia
NASA	- National Aeronautics and Space Administration
NCAR	- National Center for Atmospheric Research
NCEAS	- National Center for Ecological Analysis and Synthesis

NCEP/NCAR	- National Centers for Environment Prediction/National Center for Atmospheric Research
NDVI	- Normalized Difference Vegetation Index
NEM	- Northeast monsoon
NIMA	- National Imagery and Mapping Agency
NLAYERS	- Number of layers in the soil
NorESM1-M	- Norwegian Climate Center's Earth System Model
OND	- October, November and December
PCMDI	- Program for Climate Model Diagnosis and Intercomparison
PCPD	- Average numbers of days of precipitation for each month
PCPMM	- Average total of monthly precipitation in millimeters
PCPSKW	- Skew coefficient for daily precipitation for each month
PCPSTD	- Standard deviation for daily precipitation in each month, expressed as in mm of water per day
PDF	- Probability distribution function
PERSIANN	- Precipitation Estimation from Remotely Sensed Information Using Artificial Neural Network
PERSIANN-CDR	- Precipitation Estimation from Remotely Sensed Information Using Artificial Neural Network – Climate Data Record
PM	- Peninsular Malaysia
PMW	- Passive microwave
PR	- Precipitation radar
PR_W1	- Probability of a wet day following a dry day for each month
PR_W2	- Probability of a wet day following a wet day for each month
PRMS	- Precipitation Runoff Modelling System
PUB	- Public Utility Board
QUEST-GSI	- Quantifying and Understanding the Earth System – Global Scale Impacts
RAIN_YRS	- Number of years of maximum monthly half hour rainfall data used to define values for averages per month of the year.
RAINHHMX	- Most extreme half-hour precipitation for each month
RB	- Relative bias
RCMs	- Regional Climate Models

RCP	- Representative Concentration Pathway
RegHCM-PM	- Regional Hydroclimate Model of Peninsular Malaysia
RegHCM-SS	- Regional Hydroclimate Model of Sabah and Sarawak
ROCK1	- Rock fragment content of the first layer of soil
SAJ	- Johor Water Company
SAND1	- Sand content of the first layer of soil
SCS	- Soil Conservation Services
SDSM	- Statistical Downscaling Model
SILT1	- Silt content of the first layer of soil
SLURP	- Semi-distributed Land Use-based Runoff Processes
SNSB	- Swedish National Space Board
SOL_ALB1	- Moist soil albedo of the first layer of soil
SOL_AWC1	- Available water capacity of the first layer of soil
SOL_BD1	- Moist bulk density of the first layer of soil
SOL_CAL1	- Calcium carbonate content of the first layer of soil
SOL_CBN1	- Organic carbon content of the first layer of soil
SOL_CRK	- Crack volume potential of soil
SOL_EC1	- Soil electrical conductivity of the first layer of soil
SOL_K1	- Saturated hydraulic conductivity of the first layer of soil
SOL_PH1	- Soil PH of the first layer of soil
SOL_Z1	- Depth from soil surface to bottom of the first layer of soil
SOL_ZMX	- Maximum rooting depth of soil profile
SOLARAV	- Average daily solar radiation for each month
SPOT	- Satellite Pour l'Observation de la Terre
SPPs	- Satellite Precipitation Products
SRES	- Special Report on Emissions Scenarios
SRTM	- National Aeronautics And Space Administration Shuttle Radar Topography Mission
SSTC	- Technical and cultural services
SUFI-2	- Sequential Uncertainty Fitting
SWAT	- Soil and Water Assessment Tool
SWAT-CUP	- SWAT Calibration And Uncertainty Procedures
SWM	- Southwest monsoon

TCI	- TRMM Combined Instrument
TEXTURE	- Texture of soil layer
TMI	- TRMM Microwave Imager
TMPA	- TRMM Multi-satellite Precipitation Analysis
TMPI	- Threshold-matched precipitation index
TMPMN	- Average daily maximum air temperature for each month in degrees celsius
TMPMX	- Average daily maximum air temperature for each month in degrees celsius
TMPSTDMN	- Standard deviation for daily minimum air temperature for each month in degrees celcius
TMPSTDMX	- Standard deviation for daily maximum air temperature for each month in degrees Celsius
TOPMODEL	- Topography Based Hydrological Model
TRMM	- Tropical Rainfall Measuring Mission
UBC	- University Of British Columbia Model
UKMO	- United Kingdom Meteorological Office
UNEP	- United Nations Environment Program
UNISDR	- United Nations International Strategy for Disaster Reduction
USA	- Unite States of America
USD	- United States Dollar
USDA	- United States Department of Agriculture
USGS	- United States Geological Survey
USLE_K1	- USLE equation soil erodibility (K) factor of the first layer of soil
VIC	- Variable Infiltration Capacity
VIS	- Visible
WCRP	- World Climate Research Programme
WELEV	- Elevation of the weather station in meters above mean sea level
WGN	- Weather generator
WLATITUDE	- Latitude of the weather station
WLONGTITUDE	- Longitude of the weather station

- WMO - World Meteorological Organization
- WNDAY - Average daily wind speed for each month

LIST OF APPENDICES

APPENDIX	TITLE	PAGE
A	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 2.6 scenario from 2015 to 2034	238
B	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 2.6 scenario from 2045 to 2064	242
C	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 2.6 scenario from 2075 to 2094	246
D	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 4.5 scenario from 2015 to 2034	250
E	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 4.5 scenario from 2045 to 2064	254
F	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 4.5 scenario from 2075 to 2094	258
G	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 8.5 scenario from 2015 to 2034	262
H	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 8.5 scenario from 2045 to 2064	266

I	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 8.5 scenario from 2075 to 2094	270
J	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 2.6 scenario from 2015 to 2034	274
K	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 2.6 scenario from 2045 to 2064	278
L	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 2.6 scenario from 2075 to 2094	282
M	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 4.5 scenario from 2015 to 2034	286
N	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 4.5 scenario from 2045 to 2064	290
O	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 4.5 scenario from 2075 to 2094	294
P	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Johor River Basin under RCP 8.5 scenario from 2015 to 2034	298
Q	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 8.5 scenario from 2045 to 2064	302
R	Monthly SWAT-simulated streamflow (m^3s^{-1}) at various sub-basins in Kelantan River Basin under RCP 8.5 scenario from 2075 to 2094	306
S	Publications during the period of study for Ph.D.	310

CHAPTER 1

INTRODUCTION

1.1 Background of Study

Water is crucial for human survival. The distribution of fresh water in our planet is about 2.5% of the total amount of water, of this 68.6% is deposited in glaciers and ice caps, 30.1% in ground water and only 1.3% is found in surface water bodies such as swamps, rivers and lakes (Gleick, 1993). Surface water exists in very small amounts and is exposed to many threats caused by human activities including deforestation, pollution and climate change (e.g. Tu *et al.*, 2015; Wang *et al.*, 2015). Water resources are affected by global change, for example water demand increases due to population growth and water availability imbalance owing to climate and land use changes.

Streamflow is the flow of water in channels, rivers or streams and is one of the major components of the hydrological cycle. In general, streamflow is composed of a “slow” component called baseflow (Hall, 1968) and a “fast” component referred as surface runoff (Jakeman and Hornberger, 1993). Both components are mainly controlled by basin attributes such as soil, geology, climate, topography and land cover

type (Kirkby *et al.*, 2002; Price, 2011). Streamflow serves an important role in drinking water, hydropower, transportation, agriculture and the industrial sector.

Nevertheless, streamflow also brings some negative implications for society and environment such as flood and drought events (Hannaford, 2015). Based on the EM-DAT International Disaster Database, flood events from 1900 to 2014 have caused about 7 million deaths, with 3.6 billion people affected and 670 billion USD damage (<http://www.emdat.be/database>, accessed on 17 December 2014). Most of the severe cases occurred in tropical developing countries (UNISDR, 2011). During the December 2014 to January 2015, the monsoon flood in Malaysia affected more than 230,000 people, caused 23 deaths and cost 560 million USD damage.

Streamflow information is necessary for most of the water resources management projects (Seyam and Othman, 2015). Time series of streamflow record is crucial for flood forecasting, climate change assessment, water quality evaluation and aquatic habitat simulation (Smith and Pavelsky, 2008; Bourdin *et al.*, 2012; Ravazzani *et al.*, 2015). Better streamflow prediction and forecasting gives multiple benefits such as better water use management through anticipation of river inflow, improved capability in flood forecasting and monitoring as well as a better understanding of river water dynamics for hydropower planning (Barrett *et al.*, 2008). However, the availability and scarcity of in-situ streamflow measurements remain a critical issue in many developing countries. In many river basins, lack of data often hampers the judicious management of water resources (Hirpa *et al.*, 2013).

Remote sensing is known as a technology to retrieve information about an object without any physical contact to that object (Congalton, 2010). The ability of remote sensing in streamflow measurement has been identified and proven, either by providing input parameters to streamflow modelling or indirectly estimating the streamflow through satellite imagery (Rango, 1994; Bjerklie *et al.*, 2003; Brakenridge *et al.*, 2005; Sun *et al.*, 2015). The potential of remote sensing for providing information to hydrologists has been well recognized since the ability of Landsat-1

imagery to provide land use information for hydrological modelling (Rango, 1994; Pietroniro and Leconte, 2005).

Generally, the application of remote sensing in hydrology can be divided into three main classes including: (1) the delineation of known water features such as lakes and rivers; (2) estimating information like geological features or land use through classification and interpretation of satellite imagery; (3) the retrieval of hydrological parameters such as precipitation, evaporation, soil moisture and a digital elevation model directly from satellite imagery (Pietroniro and Leconte, 2000; Pietroniro and Prowse, 2002). Compared to the in-situ measurements by point-based gauge stations, the benefits of remote sensing to hydrology include the provision of spatial data of the area of interest, long term and global information for remote or inaccessible areas (Engman, 1996; Ritchie and Rango, 1996; Silvestro *et al.*, 2015). However, the uncertainties of satellite data in hydrological modelling and calibration of a hydrological model using satellite data are challenging topics, which are main focuses of this study.

Precipitation is an important input of water in the hydrological cycle and is the main source of catchment water (Kidd and Levizzani, 2011). Accurate and reliable precipitation information is therefore necessary to ensure better water resource management and decision-making in the various areas using water, such as agriculture, industries, and cities. Satellite precipitation products (SPPs) are widely accepted as an alternative source to overcome the limitations of ground techniques (Pan *et al.*, 2014). Recently, satellite information has become available at high spatial (up to 0.25 °) and temporal (near real time) resolutions and over large areas (near global). However, estimations using SPPs are subjected to bias and stochastic errors, which depend highly on the hydro-climatic characteristics of a region (Barrett 1993; Yilmaz *et al.*, 2005). Therefore, performance evaluation of SPPs in different regions is essential to enable users and algorithm developers to better understand and quantify such errors.

Digital Elevation Models (DEMs) are one of the most important spatial input datasets in hydrological modelling (Wu *et al.*, 2008; Bourdin *et al.*, 2012). DEMs constitute a key spatial layer for estimating a watershed's channel networks, slope gradients, flow direction and accumulations, and several other controls of the water movements in landscapes (Moore *et al.*, 1991; Wechsler, 2007). While the lack of DEM availability was an important issue a decade ago, this problem has largely been solved due to the widespread availability of global remotely sensed DEM products (Lin *et al.*, 2013) such as from the Advanced Space-borne Thermal Emission and Reflection Radiometer (ASTER) and the Shuttle Radar Topography Mission (SRTM). However, the selection of appropriate products for hydrological applications remains an important issue.

Land use change is likely to impact the water resources availability and hydrological cycle by modification in infiltration rate, interception and evapotranspiration (Chawla and Mujumdar, 2015). A multitude of natural factors such as seasonal variations in land cover characteristics and natural maturation of forest can lead to changes in land use (Öztürk *et al.*, 2013). Besides that, anthropogenic factors include deforestation, urbanization, agriculture activities, and forest management practices would also change the land use pattern. Recently, many researchers have evaluated the relationship between hydrological cycle components and land use changes (Li *et al.*, 2015; Mateus *et al.*, 2015; Nazir *et al.*, 2015; Soulis *et al.*, 2015).

Climate change is generally known as the changes in the earth's climatic condition and persists for a long term period (IPCC, 2012). The climate system is very sensitive and closely interconnected with the hydrological cycle (Zhao *et al.*, 2015). Any perturbation of the climate system will temporarily or permanently alter the hydrological cycle and lead to a significant impact on water resources (Bates *et al.*, 2008). The current water management systems are expected to be unable to cope with future climate changes; therefore, integration of climate impact information in the adaptation strategy is now a priority for each country (Wilby *et al.*, 2009).

The Fifth Assessment Report of the Intergovernmental Panel on Climate Change (IPCC-AR5) chapter three - the most up-to-date scientific assessment of past, present and potential future climate on freshwater resources - reported that water-related hazards such as river floods and droughts are expected to increase significantly with increasing precipitation and temperature (Jiménez Cisneros *et al.*, 2014). Hence, the potential of land use and climate change impacts on the hydrological cycle have received attention from numerous researchers in different regions of the world (Maurer *et al.*, 2009; Tsai and Huang, 2011; Wang *et al.*, 2012; Dessu and Melesse, 2013; Ficklin *et al.*, 2013a; Khoi and Suetsugi, 2014a; Gao *et al.*, 2015).

Generally, most of the hydro-climatic impact assessment studies adopt a “top down” framework by applying downscaled projected climate data of General Circulation Models (GCMs) into a calibrated hydrological model (Maurer *et al.*, 2009; Taye *et al.*, 2011) (Figure 1.1). Each component of the “top down” framework (Figure 1.1) contains a degree of uncertainty which leads to changes in the measured impact response (Wilby *et al.*, 2009). Hence, each component’s uncertainties need to be quantified in order to provide robust impact information (Wilby *et al.*, 2008).

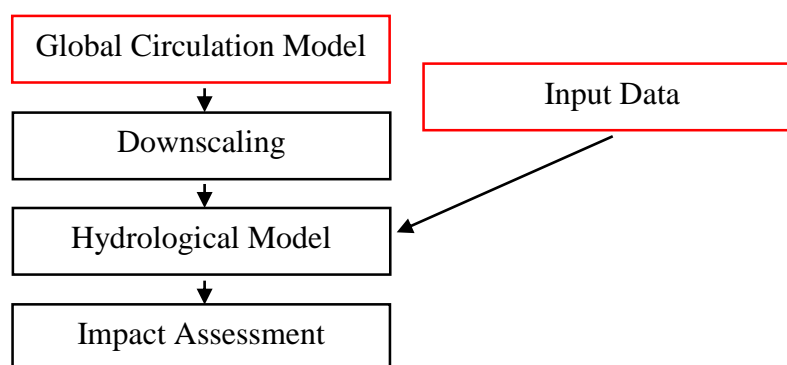


Figure 1.1 “Top down” framework for climate change impact on hydrological cycle studies (red colour boxes are components that are addressed in this study)

The selected hydrological model, the Soil and Water Assessment Tool (SWAT) is one of the most popular models. It was developed by the United States Department of Agriculture (USDA) to evaluate the impact of agriculture management on water at a basin scale (Arnold *et al.*, 1998). SWAT has proved to be a reliable tool for studying climate change impact on streamflow around the world (Zeng *et al.*, 2012; Kankam-Yeboah *et al.*, 2013; Le *et al.*, 2015; Mehdi *et al.*, 2015; Serpa *et al.*, 2015; Uniyal *et al.*, 2015). There are relatively few studies which have applied SWAT to simulate streamflow in Malaysia.

1.2 Statement of the Problem

With the development of remote sensing and computer systems, the traditional basin delineation approaches are now rapidly being replaced by automatic approaches. For example, DEM-based basin delineation gives advantages such as cost saving, efficient and faster processing. However, such an approach might lead to sub-basins and river networks mismatch issues, especially in a plain polder area (Luo *et al.*, 2011). Therefore, an improved DEM-based basin delineation and river network correction approach is required to improve modelling of streamflow with the realistic hydrological process.

The capability of a hydrological model to accurately simulate streamflow is highly dependent on the quality of the input data (Bourdin *et al.*, 2012). Satellite precipitation data potentially constitutes an alternative to sparse rain gauge networks for assessing the spatial distribution of precipitation (Casse *et al.*, 2015). However, the uncertainties in the satellite precipitation might propagate in hydrological models and need to be accounted for before applying it to hydrological models (Nikolopoulos *et al.*, 2010; Bajracharya *et al.*, 2015; Skinner *et al.*, 2015). Hence, evaluation of SPPs performance in rainfall estimation before hydrological modelling is very important to reduce uncertainties in final output.

Studies about the evaluation of SPPs in Malaysia appear to be limited. Varikoden *et al.* (2010) evaluated the daily precipitation from TRMM 3B42V6 data in Peninsular Malaysia, which covers only about 40% of the total area of Malaysia, using for validation only four precipitation gauges. Semire *et al.* (2012) validated the TRMM Microwave Imager (TMI) 2A12, 3B42V6, 3B43V6, and Global Precipitation Climatology Center (GPCC) with the monthly precipitation data collected over 10 years (2001–2010) from 22 precipitation gauges distributed over Malaysia. Both studies showed that 3B43V6 performs well over Malaysia, with a $\pm 15\%$ error bias at monthly scale. However, these studies have compared only one or two SPPs, thus limiting their conclusions. There are several other available SPPs that need to be tested such as the Global Precipitation Climatology Project (GPCP), Precipitation Estimation from Remotely Sensed Information Using Artificial Neural Network (PERSIANN), and Climate Prediction Center Morphing Technique (CMORPH). Moreover, the accuracy assessment of these products is also required at daily time steps.

The climate projection uncertainty assessment is a quantification of the feasible threshold of future hydro-climatic conditions (Dessu and Melesse, 2013). The Quantifying and Understanding the Earth System – Global Scale Impacts (QUEST-GSI) methodology developed by Todd *et al.* (2011) is an advanced method involving the unification of climate impacts and uncertainties scenarios for better comparison between different study areas and applying prescribed warming scenarios to advice mitigation policy. However, the QUEST-GSI methodology applied previous GCMs from the phase 3 Coupled Model Intercomparison Project (CMIP3) and emission scenarios from the Special Report on Emission Scenarios (SRES). Hence, modification of the QUEST-GSI methodology is essential to fit with the latest climate research development.

The regional scale hydro-climatic assessment studies especially on the water sector in Malaysia are still limited, and there is no related research listed in the Intergovernmental Panel on Climate Change (IPCC) Fourth Assessment Report (AR4) (IPCC 2007; Tangang *et al.*, 2012). Hydro-climatic impacts research in Malaysia has been receiving much more attention since 2007, due to the increase in climate impact

awareness. Shaaban *et al.* (2012) have evaluated impacts of climate change on several major basins of Peninsular Malaysia for the 2025 to 2034 and 2041 to 2050 periods by using the Canadian Centre for Climate Modelling and Analysis (CCCma) under the IS92a emission scenario. However, the application of the latest emission scenarios (Moss *et al.*, 2010) and multiple GCMs assessment (Palmer *et al.*, 2008) is very important to reflect more robust impact analysis. In this study, the latest CMIP5 GCMs and emission scenarios released by the IPCC in 2014 were used, and the missing “end of the century” assessment was included.

Most of the studies were concerned with the changes of a single factor, i.e. either climate changes (Shaaban *et al.* 2011) or land-use changes (Amini *et al.* 2011). Adnan and Atkinson (2011) considered both impacts of land-use and climate changes in Malaysia by using a statistical analysis approach. The study was conducted in northeast Peninsular Malaysia and they found an increasing trend of streamflow and precipitation during the wet season. However, no hydrological modelling impact studies that consider both factors (land-use and climate) have been performed in other part of Peninsular Malaysia. This study aims to fill this research gap.

There are only a few studies that applied SWAT to simulate hydrological processes in Malaysia. Memarian *et al.* (2014) applied SWAT in the Hulu Langat River Basin to assess impact of land-use change and reported the Nash-Sutcliffe Efficiency (NSE) value of 0.73 in validation period. In the Upper Bernam River Basin, Alansi *et al.* (2009) reported the validated SWAT’s NSE values range from 0.68 to 0.78. These studies found that SWAT showed moderate to good performance in northern and western Peninsular Malaysia. However, the capability of the SWAT model to simulate streamflow in east-coast and southern Peninsular Malaysia is still unclear. Parameters in the empirical equations within SWAT were derived from climate condition in the USA (Neitsch *et al.*, 2011) and these should be modified for application to tropical conditions, in order to obtain better hydrological simulation.

1.3 Research Questions

1. Which digital elevation model is suitable for SWAT modelling?
2. Which satellite precipitation product is suitable for providing rainfall information for hydro-climatic assessment?
3. What is the performance of the SWAT model in Malaysia?
4. To what extent were land use and climate changes influencing the streamflow?
5. How much change will there be in future streamflow under climate change impacts?

1.4 Objectives of the Study

The aim of this study is to explore and quantify the variability of climate change on streamflow variability. The specific objectives are outlined as follows:

1. To quantify the uncertainties of digital elevation models.
2. To evaluate and reduce the uncertainties of satellite precipitation products.
3. To modify the Soil and Water Assessment Tool (SWAT) for streamflow simulation in Malaysia.
4. To investigate the impacts of historical land-use and climate change on streamflow.
5. To assess the impacts and uncertainties of future climate change on streamflow.

1.5 Scope of the Study

The scope of this study can be divided into two parts. The first part involves the development of a SWAT model for the JRB. The input data sets are DEM, SPP, observed climate data (1985 to 2012), observed streamflow data (1985 to 2012), land use map and soil map. The SPP, land use map and DEM were generated from satellite data, while the observed climate, streamflow and soil map were collected from the relevant government agencies. The Latin hypercube sampling by one at a time design (LH-OAT) developed by Van Griensven *et al.* (2006) was used to perform sensitivity analysis and calibration procedure. It is an automatic sensitivity analysis embedded in the SWAT-CUP tool.

The second part was focused on the climate projection uncertainty analysis by using the modified QUEST-GSI methodology. The methodology was modified by incorporating the latest GCMs from the Phase 5 Coupled Model Intercomparison Project (CMIP5) and the Representative Concentration Pathways (RCPs) 2.6, 4.5 and 8.5 which represent low, medium and high emission scenarios, respectively.

Five GCMs with a good capability of historical precipitation simulation were selected from 18 GCMs for an average ensemble process and these GCMs were used for the hydro-climatic assessment. The ensemble of GCMs was used to replace a single GCM in the QUEST-GSI methodology. The periods set for impact assessment are 1985-2004(baseline), 2015-2034 (beginning of the century), 2045-2064 (mid of the century) and 2075-2094 (end of the century). The developed climate scenarios were used as inputs into the calibrated SWAT to assess impacts and uncertainties of climate change on future streamflow.

1.6 Significance of the Study

In hydrology, the DEM plays a vital role in providing the basin's topography information including channel network, slope and flow direction. Due to rigorous spatial data policies in most countries, many researchers prefer to choose public domain DEM products. Nowadays, there are many satellite sensors capable of producing DEMs and this will lead to difficulties for hydrologists to choose an appropriate DEM for their work. This study can be used by hydrologists to choose the best DEM for their study prior to the development of the hydrological modelling. Besides that, an improved DEM-based basin delineation and river network correction approach was introduced to correct the river network formation in SWAT model.

Precipitation is an important input of water in the hydrological cycle; hence, accurate and reliable precipitation information is therefore necessary to ensure better water resources management and decision-making. SPPs are widely accepted as an alternative source to overcome the limitation of ground techniques. This study presents a more comprehensive evaluation across Malaysia than was previously available in term of the number of gauge stations, evaluated satellite products, and the evaluated aspects (rainfall intensity analysis and flood event assessment). Moreover, this study presents the first attempt to evaluate the uncertainty and rain-detection ability of each of the SPPs over Malaysia. This study can be a guide on the selection of better precipitation products and bias correction of the selected product over Malaysia for hydrological modelling.

Land use maps were derived from multi-temporal satellite sensors images, and then applied into the SWAT model to evaluate the impacts of past land-use and climate change on streamflow of the studied basins. The knowledge on how land-use change on basin hydrology is important for local authority and policy makers to formulate and implement effective strategies to minimize the undesirable effects of future land-use change.

Application of the SWAT model in hydrological cycle simulation is still limited in Malaysia. This study presents the first attempt to evaluate the applicability of SWAT in the Johor River Basin (JRB) and Kelantan River Basin (KRB). The modified SWAT model can reasonably represent better streamflow simulation of the studied basins in Malaysia. Besides that, the SWAT's calibrated parameters of both basins can act as a reference for researchers or water resources managers to calibrate the SWAT model in tropical regions.

Building on current scientific knowledge, this study investigates the potential of climate changes on streamflow characteristics, and the associated uncertainties in the climate projection. The modified QUEST-GSI methodology is the most comprehensive and up-to-date climate scenario that is designed particularly for the quantification of climate projection uncertainty on streamflow simulation. This methodology has contributed to new theoretical and knowledge on how to better quantify the future climate projection uncertainties. This methodology is not only applicable for the evaluation on hydrological modelling, but also other applications such as environmental and ecological modelling.

The hydrological cycle varies regionally and this study contributes to factual information about the local hydrological cycle. Fresh water is one of the most important elements for human survival. Large quantities of water are essential for agriculture, industrial and hydropower purposes. As streamflow is a measure of water availability (Milly *et al.*, 2005), so the impact of climate change on future streamflow is vital for policy makers and water managers to develop better adaptation planning in terms of formulation of policies for investments in irrigation systems, agriculture, hydropower production and flood protection.

1.7 Structure of Thesis

This thesis consists of six chapters. The first chapter presents the background of the study, statement of the problem, study objectives, scope of the study and the structure of thesis. Chapter two presents a review of the hydrological cycle, hydrological modelling, remote sensing in hydrological modelling, DEM uncertainty assessment, satellite precipitation uncertainty assessment, land-use change and hydrological modelling, climate change and hydrological modelling, GCMs, downscaling of GCMs, uncertainties of climate change, studies of climate change impacts on streamflow in Asia and Malaysia.

Chapter three presents the research methodology for this study, the selected study area, essential data sets, satellite data processing and validation, SWAT model modification and set up, historical land-use and climate change analysis, modification of the QUEST-GSI methodology and application of the modified QUEST-GSI methodology in the calibrated SWAT.

Chapter four presents the main results of DEMs and SPPs validation, river network correction and bias correction of the selected satellite precipitation product. In chapter five, the main findings such as SWAT calibration and validation, GCMs performance evaluation, impacts and uncertainties of climate change on streamflow of the JRB and KRB are presented. Finally, the summary, conclusion and future work recommendations of the study are described in chapter six.

to simulate more the tropical hydrological cycle reliably, in order to gain better simulation results.

- (d) More climate scenarios should be established to understand better the range of climate impacts on streamflow. In addition, the modified QUEST-GSI methodology should be applied to other regions to check its capability.
- (e) A better ensemble of GCMs approach should be developed in order to evaluate potential climate impact on streamflow more effectively. For example, applying the weight factors to GCMs based on their performance in simulating historical climate variables.

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